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# Water stable isotopes reveal groundwater vulnerability to land use fragmentation and climate variability in central Honduras

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## Manuscript Components

**Title:** 17 words

**Abstract:** 222 words

**Main text:** 3,797 words (excluding references)

**Supplementary Tables:** 2 (tables caption text: 49 words; online repository)

**Figures:** 8 in color (figures caption text: 423 words)

**Total references:** 60

**Running title:** *Groundwater vulnerability in the tropics*

## Abstract

The inter-mountainous region of central Honduras has been experiencing abrupt urban water shortages during the last decade. Land use fragmentation to increase pasture, crop, and peri-urban areas has rapidly reduced surface water quantity and quality. Here we present a 3-yr (2018-2020) tracer study within the headwaters of the Choluteca River basin (2,949 km<sup>2</sup>). We sampled rainfall (weekly N=156; daily N=270), drilled wells (N=166; up to 300 m depth), boreholes (N=70; 4-12 m depth), and springs (N=128) to assess the spatio-temporal connectivity between rainfall and groundwater recharge elevations (MREs). Rainfall isotopic seasonality from the dry to the wet season is characterized by a clear W-shaped pattern. HYSPLIT trajectory analysis reveals three main moisture sources: 73% (east, Caribbean Sea), 17% (southwest, Pacific Ocean), and 10% (north; Gulf of Mexico). Groundwater sources exhibit a strong meteoric origin with a slight evidence of secondary evaporation. MREs range between 813 to 2,130 m asl with a mean value of ~1,600 m asl. Seasonal isotopic variability along with the influence of rapid infiltration limited the MRE method performance as follows: springs (60%), borehole (77%), and drilled wells (93%). Isotope-informed MREs coincide in large degree with coniferous forest, pasture, and crop areas. These results are intended to guide the mapping and delineation of critical recharge areas in central Honduras to enhance water quality and quantity during dry periods.

**Keywords:** Honduras; water stable isotopes; rainfall; groundwater connectivity; critical recharge zones.

49 **1. Introduction**

50 Honduras has some of the highest poverty rates in the Latin America and  
51 Caribbean region ([Mateo, 2021](#)). Out of a total population of 9,523,621 inhabitants (INE,  
52 2021), an estimated 48.3% and 22.9% lived below the national poverty and extreme  
53 poverty lines in 2018, respectively ([World Bank Group, 2020](#)). It is expected that  
54 COVID-19 pandemic and the direct impacts of hurricanes ETA and IOTA may  
55 exacerbated the social inequality in a near future ([World Bank Group, 2020](#); [Zúñiga-](#)  
56 [Moya et al., 2020](#); [Durón et al., 2021](#); [Zambrano et al., 2021](#)).

57 Water availability (in quality and quantity) is one of the most challenging public  
58 services in Honduras. Although nearly two thirds of Honduras have access to drinking  
59 water, 90% of systems provide only intermittent services (e.g., few hours weekly or  
60 biweekly) ([Balthasar and Zairis, 2011](#); [Grillos et al., 2021](#)). As a common denominator  
61 people have reported being forced to purchase expensive bottled or trucked water  
62 during the prolonged dry season (5-7 months, from November to May) ([Grillos et al.,](#)  
63 [2021](#); [Rodríguez-Cruz et al., 2021](#)).

64 The central region of Tegucigalpa (i.e., Francisco Morazán Department) hosts  
65 nearly 18% of the country's population with a relatively high population density  
66 (198.4 inhabitants /km<sup>2</sup>) ([INE, 2021](#)). This region is characterized by a highly urbanized  
67 intra-mountain valley including recent urban settlements across high elevation peri-  
68 urban areas. In the past decades, pine bark beetle has triggered the major loss of  
69 coniferous forests in central Honduras, decreasing thousands of hectares of pine trees  
70 and mixed forests ([Valdez et al., 2017](#)). The latter has concatenated a rapid change in  
71 land use with an increase in pasture and crops areas which in turn has rapidly reduced  
72 surface water quantity and quality ([Lee, 2000](#); [Vignola et al., 2015](#); [Blair et al., 2019](#);  
73 [Gotlieb et al., 2021](#)). In addition, prolonged El Niño-induced droughts in Central America  
74 have reduced the availability of water-in particular for the western inter-mountainous  
75 region of the Dry Corridor of Honduras ([Vignola et al., 2015a](#); [Quesada-Hernández et](#)  
76 [al., 2019](#); [Depsky et al., 2020](#); [Olivera et al., 2021](#); [Rodríguez-Cruz et al., 2021](#)). Rainfall  
77 deficits are recurrently affecting the volume of the main surface water reservoirs in  
78 central Honduras.

79           The combination of these factors is increasing illegal groundwater exploitation  
80 without a proper scientific baseline regarding the timing, magnitude, and identification of  
81 critical recharge areas in central Honduras. Therefore, the main goal of this study was  
82 to a) determine the spatio-temporal water isotopes variability (in rainfall, surface, and  
83 groundwater) and dominant atmospheric moisture sources and to b) understand the  
84 governing recharge mechanisms and mean recharge elevations (MREs) in the  
85 headwaters of the Choluteca River basin (in central Honduras). We hypothesized that  
86 groundwater recharge is governed by infiltration within the high elevations of the basin  
87 and in less degree by urban recharge. Our results highlight basin regions where water  
88 infiltration is affected by land use fragmentation as well as critical headwater recharge  
89 areas for water protection and conservation under a changing climate.

## 90 **2. Study Area**

91           The study area comprises the headwaters of the Choluteca River basin, located  
92 in the south-central region of Honduras with a total area of 2,949 km<sup>2</sup> (Fig. 1). The basin  
93 can be divided into four main sub-basins: Choluteca Alta (1,255 km<sup>2</sup>), Guacerique-  
94 Grande (723 km<sup>2</sup>), Río del Hombre (450 km<sup>2</sup>), and Yegüare (524 km<sup>2</sup>). The elevation  
95 gradient ranges from 541 up to 2,335 m asl.

### 96 **2.1. Climate generalities**

97           Precipitation in Honduras is mainly influenced by the trade winds (E-W prevailing  
98 direction), the Intertropical Convergence Zone (ITCZ), easterly tropical waves, cold front  
99 outbreaks, troughs (i.e., low pressure systems), indirect and direct effect of tropical  
100 cyclones, and El Niño Southern Oscillation (ENSO) (Westerberg et al., 2009; [Alfaro et al., 2018](#)).  
101 However, a robust moisture source-sink analysis is still lacking. Figure 2  
102 shows a representative long-term seasonal pattern between the highlands (La Brea  
103 station; 1,610 m asl) and the Tegucigalpa valley (UNAH station; 1,063 m asl) within the  
104 headwaters of the Choluteca River basin (Fig. 1). The rainfall regime exhibits a common  
105 bimodal pattern for the Central America region ([Durán-Quesada et al., 2017](#); [Gouirand et al., 2020](#)).  
106 Monthly mean rainfall oscillates from 4.6 and 167.4 mm at the urban  
107 lowlands (UNAH station) and from 9.0 and 285.0 mm at the highlands (Fig. 2).

108           The rainy season spans from May to November. Between May and June, rainfall  
109 is mainly caused by the entry of Pacific winds which are enhanced by a) the thermal

110 gradient between the Pacific Ocean and the continental area of Central American  
111 territory and the confluence with the trade winds and b) the migration of the first tropical  
112 waves of the year from the Caribbean Sea. In general, a rainfall decrease between July  
113 and August is observed as a result of the Mid-summer drought (MSD) ([Magaña et al.,](#)  
114 [1999](#); [Small et al., 2007](#)). The ITCZ influence and the indirect/direct effect of tropical  
115 cyclones results in a notable rainfall increase during September and November.

116 During ENSO's warm phase a decrease in precipitation is observed during most  
117 of the year. Extreme dry spells in the region results in a net rainfall reduction of up to  
118 31%, which is equivalent to about 368 mm of the mean annual precipitation, while  
119 during extreme rainy events, annual precipitation can increase by 35% (corresponding  
120 to approximately 434 mm). Mean monthly temperature ranges from 17.9°C to 25.8°C,  
121 for the rainy and dry seasons, respectively. The mean annual temperature is 22°C,  
122 reaching maximum values between March and May. Mean minimum temperature are  
123 observed between December and February, ranging from 17.9°C and 18.4°C.

## 124 **2.2. Lithological outlines**

125 The basin is defined by three groups known as Honduras Group, Valle de  
126 Ángeles Group and Padre Miguel Group, including three formations: Río Chiquito,  
127 Villanueva and Matagalpa. In the study area, the main outcrops correspond to Valle de  
128 Ángeles Group and Padre Miguel Group (Fig. 3). Valle de Ángeles Group ([Rogers and](#)  
129 [O'Conner, 1993](#); [Rogers et al., 2007](#)) comprises the Río Chiquito Formation, which  
130 mainly consist of red layers of fine texture; being strata containing shales, limonites,  
131 sandstones and some layers of conglomerate of quartz; with a thickness ranging from  
132 400 to 800 meters. On this unit of sedimentary rocks lie some thin layers of volcanic  
133 rocks known as the Matagalpa Formation ([Williams & McBirney, 1969](#)), which can be  
134 described as calc-alkaline lava flows with andesitic and basaltic compositions and minor  
135 pyroclastic layers ([Barberi et al., 2013](#)).

136 The Padre Miguel Group covers an area of approximately 70,000 km<sup>2</sup> in  
137 Honduras and is well exposed in the headwaters of the Choluteca River basin (Fig. 3)  
138 ([Garza et al., 2012](#)). This group is described as a silicic volcanic sequence (ca. ~2,000  
139 m thickness) that forms part of the Central American Miocene volcanic arc built on the  
140 Chortis continental fragment ([McBirney & Williams, 1969](#); [Garza et al., 2012](#); [Barberi et](#)

141 [al., 2013](#)). The rocks consist of white, red, pink and green rhyolitic, dacitic and andesitic  
142 tuffs with quartz, feldspars, sanidine, biotite and lithic crystals of pumice and clasts from  
143 the Valle de Ángeles Group and Matagalpa Formation. They also include some  
144 sedimentary rocks of volcanic clasts and well stratified ashlar type tuffs. After the  
145 deposition of the main ignimbrite rocks, there was a time of relative calm where the  
146 environmental conditions allowed the rocks to be altered and weakened, which  
147 originated the so-called Lahars. The outcrops of laharc conglomerate materials that are  
148 widely distributed in the Tegucigalpa quadrangle consist of highly angular, subangular  
149 and rounded clasts of coarse gravel from runoffs, tuffs and clasts from the Valle de  
150 Ángeles Group. The deposit is not well sorted and appears chaotically deposited in a  
151 paleo-valley in the Valle de Ángeles Group ([McBirney & Williams, 1969](#); [Garza et al.,  
152 2012](#); [Barberi et al., 2013](#); [Braun et al., 2019](#)).

### 153 **3. Methods and Materials**

#### 154 **3.1. Rainfall collection**

155 We monitored rainfall for stable isotopes analysis using two passive Palmex  
156 collectors (weekly collection at Cerro de Hula and El Picacho; N=156) and one in-house  
157 passive collector (daily collection at UNAH; N=270) with collection funnels from 13.5 cm  
158 to 18.0 cm and high-density polyethylene (HDPE) containers of 2.3 to 3 L, respectively  
159 ([Gröning et al., 2012](#)) (Fig. 1). Rainfall collectors were installed considering the  
160 elevation gradient, rainfall moisture sources, and access for maintenance and protection  
161 of the collectors throughout the hydrological year (Supplementary Table S1; [García-  
162 Santos et al., 2021](#)). It is important to highlight that site accessibility and security  
163 conditions within the headwaters of Choluteca River basin impeded rainfall collection at  
164 higher elevations (i.e., >1,700 m asl).

#### 165 **3.2. Springs, drilled wells, and boreholes sample collection**

166 We collected samples from groundwater sources during three dry seasons  
167 (2018, 2019, and 2020) and during the MSD of 2018. Samples were collected targeting  
168 baseflow periods to better represent mean annual recharge and discharge isotopic  
169 values ([Sánchez-Murillo et al., 2016](#); [Sánchez-Murillo et al., 2020 a,b](#)). Samples were  
170 classified as drilled wells (N=166) with depths up to 300 m, boreholes (N =70) with  
171 depths between 4 and 12 m and springs (N=128). In total, we sampled 364 water

172 sources with a re-sampling rate per source of 29% for springs, 36% for boreholes, and  
173 52% for drilled wells (two or three samples per site). All the samples were collected in  
174 HDPE bottles (15 to 30 mL) and stored at 4°C to avoid fractionation (Supplementary  
175 Table S2; [García-Santos et al., 2021](#)).

### 176 **3.3. Stable isotopes analysis**

177 Stable isotopes analyses were conducted at the Stable Isotopes Research Group  
178 facilities of Universidad Nacional (Heredia, Costa Rica) using a LWIA-45-EP water  
179 isotope analyzer (Los Gatos, USA). Calibrated secondary standards were used to  
180 normalize the results as well as to assess quality and drift control procedures.  $^{18}\text{O}/^{16}\text{O}$   
181 and  $^2\text{H}/^1\text{H}$  ratios are presented in delta notation  $\delta$  (‰), with reference to the VSMOW-  
182 SLAP scale. The long-term analytical uncertainty was  $\pm 0.5$  ‰ for  $\delta^2\text{H}$  and  $\pm 0.1$  ‰ for  
183  $\delta^{18}\text{O}$ . Deuterium excess was calculated as  $d$ -excess= $\delta^2\text{H}-8\cdot\delta^{18}\text{O}$  (Dansgaard, 1954).

### 184 **3.4. HYSPLIT air mass back trajectories**

185 In Central America, moisture transport often acts as an 'atmospheric bridge' that  
186 connects the semi-closed Caribbean Sea basin and the eastern Pacific warm pool  
187 ([Sánchez-Murillo et al., 2020a,b](#)). Inter-basin moisture transport is largely modulated by  
188 the ITCZ, easterly tropical waves, sea surface temperature (SST) variation, and the  
189 indirect/direct effect of tropical cyclones. Such interactions exert a large control on  
190 rainfall seasonality and distribution ([Pan et al., 2018](#)) with variability influenced by  
191 ENSO at interannual scales ([Yun et al., 2021](#)). In this context, moisture transport was  
192 evaluated using the HYSPLIT Lagrangian model ([Stein et al., 2015](#); [Rolph et al., 2017](#)).  
193 This model uses a three-dimensional Lagrangian air mass vertical velocity algorithm to  
194 determine the average position of the air mass, which is reported at an hourly time  
195 resolution over the trajectory ([Soderberg et al., 2013](#)). Air parcel trajectories (start time  
196 12:00 p.m. UTC) were modeled 48 hours backwards in time based on the proximity of  
197 the Caribbean Sea and the Pacific Ocean. In total, 270 trajectories (UNAH daily station,  
198 Fig.1 and Table S1) were created and further divided by a) seasons (dry: December to  
199 April; wet: May to November), b)  $\delta^{18}\text{O}$  compositions and c)  $d$ -excess variability.

### 200 **3.6. Mean recharge elevations (MREs) and rainfall-groundwater connectivity**

201 Critical recharge areas within the headwaters of the Choluteca River basin were  
202 evaluated using an isotopic lapse rate under the assumption that groundwater isotope

203 ratios under baseflow conditions are representative of mean annual recharge conditions  
204 during the wet season ([Yamanaka and Tamada, 2017](#); [Sánchez-Murillo et al., 2020a,b](#)).  
205 The latter is supported by the amount of rainfall generated in the wet season with  
206 respect to the annual volume (UNAH=98%, Cerro de Hula=97%). Mean annual  $\delta^{18}\text{O}$   
207 compositions were calculated in four stations within the study area: UNAH (-5.50‰), El  
208 Picacho (-5.41‰), Cerro de Hula (-7.13‰), and Rincón de Dolores (archive IAEA data;  
209 longitude=-87.3997, latitude=14.1401) (-7.42‰), resulting in a significant ( $r^2=0.82$ ,  
210  $p<0.01$ ) linear regression between 1,063 and 1,643 m asl. A Kruskal-Wallis non-  
211 parametric ([Kruskal and Wallis, 1952](#)) one way analysis of variance on ranks was  
212 conducted to diagnose significant differences among the four sites. The differences in  
213 the median values among the sites are greater than would be expected by chance;  
214 thus, there is a statistically significant difference ( $p<0.001$ ).

215 To contextualize the relevance for water resources management of the isotope-  
216 inferred MREs within the headwaters of the Choluteca River basin, the most recent land  
217 use map (2019) (resolution 10x10 m based on satellite images from the Copernicus  
218 Sentinel-2 mission) (<https://sentinel.esa.int/web/sentinel/missions/sentinel-2>) was  
219 obtained from the National Institute of Conservation and Forestry Development  
220 (<http://www.geoportal.icf.gob.hn/geoportal/main>). Land use includes 19 classifications  
221 which denote the high degree of land fragmentation within the basin. Statistical and  
222 graphical analysis was performed using the open source statistical R language and  
223 packages (R Development Core Team, 2021). All maps were developed in ArcGIS 10.5  
224 (ESRI, USA).

## 225 **4. Results**

### 226 **4.1. Rainfall isotopic variability**

227 At the urban lowlands (UNAH station)  $\delta^{18}\text{O}$  ranged from -18.67 to +2.13‰, with a  
228 mean of -5.11‰; while  $\delta^2\text{H}$  varied from -136.2 to +19.8‰, with a mean of -32.5‰. At El  
229 Picacho station  $\delta^{18}\text{O}$  ranged from -14.91 to +0.39‰ and  $\delta^2\text{H}$  varied from -109.1 to  
230 +14.6‰; while at El Cerro de Hula station  $\delta^{18}\text{O}$  ranged from -20.32 to -0.13‰ and  $\delta^2\text{H}$   
231 values varied from -148.0 to +15.8‰ (Fig. 4A). Climate seasonality from the dry  
232 (December-April) to the wet season (May-November) is characterized by a W-shaped  
233 isotopic pattern in rainfall (Fig. 4A). The latter is consistent with the intra-seasonal

234 variability of Central America rainfall, that typically result in two or three depleted  
235 excursions during the wet season and two enriched pulses during the strongest trade  
236 winds period and the MSD. A similar pattern has been reported in Costa Rica,  
237 Nicaragua, Veracruz (México), and Guatemala ([Wassenaar et al., 2009](#); [Lachniet et al.,  
2009](#); [Pérez-Quezadas et al., 2015](#); [Sánchez-Murillo and Birkel, 2016](#); [Sánchez-Murillo  
et al., 2019](#); [Sánchez-Murillo et al., 2020 a,b](#)).

240 Figure 4B shows the temporal *d*-excess variation in rainfall across the  
241 headwaters of Choluteca River basin. Relative high *d*-excess values (>20‰), at El  
242 Cerro de Hula and El Picacho stations during dry season months (December to April),  
243 indicate moisture recycling and the influence of cold front events from the Gulf of  
244 Mexico. After the onset of the rainy season in mid-May, mean *d*-excess varied between  
245 10-15‰ in both high elevation stations (Fig. 4B). However, at the urban lowland *d*-  
246 excess values were consistently below 10‰ mainly due to the combination of strong  
247 sub-cloud evaporation and high ambient temperatures. Isotope ratios resulted in a  
248 highly significant meteoric water line (CBMWL):  $\delta^2\text{H}=7.77 \cdot \delta^{18}\text{O}+9.14$  ( $\text{Adj.}r^2=0.97$ ;  
249  $p<0.001$ ;  $N=426$ ) (Fig. 5), in agreement with MWLs within the Pacific slope of Central  
250 America ([Sánchez-Murillo et al., 2020a](#)).

#### 251 **4.2. Springs, drilled wells, and boreholes isotopic variability**

252 Figure 5 shows a combination of dual-isotope diagrams for springs, drilled wells  
253 and boreholes. Overall, all sources exhibited a strong meteoric origin with slight  
254 evidences of secondary evaporation. It is important to remark that mean  $\delta^{18}\text{O}$  for all  
255 water sources is within the mean annual rainfall  $\delta^{18}\text{O}$  compositions: UNAH (-5.50‰), El  
256 Picacho (-5.41‰), Cerro de Hula (-7.13‰). In the drilled wells,  $\delta^{18}\text{O}$  ranged from -9.18  
257 to -3.40‰, with a mean of  $-7.05 \pm 0.69\text{‰}$  ( $1\sigma$ ) (Fig. 5A). At the boreholes,  $\delta^{18}\text{O}$  values  
258 fluctuated between -8.71 and -3.76‰, with a mean value of  $-6.52 \pm 1.02\text{‰}$  ( $1\sigma$ ) (Fig. 5B).  
259 Across the spring sites,  $\delta^{18}\text{O}$  ranged from -9.11‰ to -1.14‰, with a mean of -  
260  $7.08 \pm 1.02\text{‰}$  ( $1\sigma$ ) (Fig. 5C).

#### 261 **4.3. HYSPLIT air mass back trajectories**

262 From 270 air parcel trajectories, 93% ( $N=250$ ) and 7% ( $N=20$ ) corresponded to  
263 the wet and dry seasons, respectively. Based on the main trajectory direction  
264 (quadrangle centered on Tegucigalpa), moisture transport origin can be divided as

265 follows: 73% (east; Caribbean Sea), 17% (southwest; Pacific Ocean), and 10% (north;  
266 Gulf of Mexico). In the case of  $\delta^{18}\text{O}$ , enriched events between -2.13 and -6.26‰  
267 originated mostly from the Caribbean Sea with some early season events coming from  
268 the eastern Pacific Ocean (mainly in April). Depleted events (up to -18.67‰ in  $\delta^{18}\text{O}$ )  
269 occurred mainly during the second leg of the rainy season and were originated from the  
270 southwest and in less degree from the Caribbean basin during the activation of the  
271 cyclonic activity (Fig. 6A). Cold front events from the Gulf of Mexico often result in  
272 enriched isotopic light rainfalls ([Sánchez-Murillo et al., 2019](#); [Welsh and Sánchez-](#)  
273 [Murillo, 2020](#)).

274 Similarly, high *d*-excess rainfall events (>12.64‰) occurred mainly from the  
275 northeastern and southwest regions of the Caribbean Sea and Pacific Ocean,  
276 respectively. Rainfall events near the mean global *d*-excess value (10‰) were  
277 originated mainly from the Caribbean Sea and in less degree from the Pacific Ocean  
278 (Fonseca Gulf). Small rainfall events with relative low *d*-excess were indistinctive  
279 generated from both oceanic basins (Fig. 6B) and are related to sub-cloud evaporation  
280 in the Tegucigalpa valley ([Managave et al., 2016](#); [Salamalikis et al., 2016](#); [Xia and](#)  
281 [Winnick, 2021](#)).

#### 282 **4.4. Mean recharge elevations and rainfall-groundwater connectivity**

283 The isotopic lapse rate resulted in a significant ( $r^2 = 0.82$ ;  $p < 0.01$ ) slope of -3.6‰  
284 per km of elevation between 1,063 and 1,643 m asl within the headwaters of the  
285 Choluteca River basin. Recently [Sánchez-Murillo et al. \(2020b\)](#) reported a pantropical  
286 spectrum of isotopic lapse rate ranging from -3.5 to -0.5‰/km, with a pantropical mean  
287 of -2.2‰/km. Similarly, [Poage and Chamberlain \(2001\)](#) reported a global isotopic lapse  
288 rate of -2.8‰/km. In the study area, MREs ranged between 813 to 2,130 m asl (Fig. 7A)  
289 with a mean value near 1,600 m asl (Fig. 7B). The seasonal isotopic variability, and  
290 consequently, the impact of near surface infiltration from recent rainfall events limited  
291 the inverse recharge elevation calculation. The latter is more evidence in springs 40%  
292 (51 sources of 128) and boreholes 23% (16 sources of 70); while the method exhibited  
293 a lower rate of invalid calculations in drilled wells with only 7% (12 sources of 166) (Fig.  
294 7A). This result is also consistent with the pattern of water sources along the CBMWL  
295 with a greater dispersion in the springs and boreholes than drilled wells (Fig. 5). Drilled-

296 wells exhibited a greater damping effect than shallow springs and boreholes ([Berman et](#)  
297 [al., 2009](#); [Dehaspe et al., 2018](#); [Hu et al., 2020](#)).

## 298 **5. Discussion**

### 299 **5.1. Land use, groundwater recharge processes and critical recharge elevations**

300 Isotope-inferred MREs were divided in five classes within the headwaters of the  
301 Choluteca River basin, as follows: (a) 541 to 1,090 m asl, (b) 1,091 to 1,470 m asl, (c)  
302 1,471 to 1,678 m asl, (d) 1,679 to 1,854 m asl, and (e) 1,855 to 2,335 m asl. These  
303 classes cover 42%, 38%, 13%, 5%, and 3% of the study area, respectively. Class (a)  
304 represents a region with a large prevalence of anthropogenic activities and soil  
305 degradation such as pastures and crops (21%), technified agriculture (8%), and urban  
306 areas (6%). This class (a) involves one of the most human-altered regions of central  
307 Honduras ([Vignola et al., 2015](#)); however, MREs indicate a recharge bias towards  
308 higher recharge elevations and more depleted events in the basin ([Jasechko and](#)  
309 [Taylor, 2015](#)). The latter is consistent with a previous study across the Pacific slope of  
310 the Dry Corridor of Central America (including Honduras), where groundwater recharge  
311 was primordially reported in highland areas (72.3%) versus rapid (13.1%) and localized  
312 (14.5) recharge ([Sánchez-Murillo et al., 2020a](#)).

313 Class (b) comprises mainly pastures and crops (15.95%) and coniferous forests  
314 affected by the pine bark beetle plague (15.22%) ([Valdez et al., 2017](#)). In contrast to  
315 class (a), this class showed a high incidence of local recharge. Although most of the  
316 drilled wells are located in the lowlands (classes and b), the most prominent isotope-  
317 inferred MREs were found between 1,471 and 1,854 m asl, corresponding to classes (c)  
318 and (d). This highlights the large disproportionality existing between MREs and  
319 headwaters dependent systems in tropical regions ([Yamanaka and Tamada, 2017](#);  
320 [Immerzeel et al., 2020](#); [Sánchez-Murillo et al., 2020 a,b](#)). Class (c) is mainly covered by  
321 pine forests affected by the bark beetle plague (19%), pastures and crops (12%),  
322 whereas in class (d) the land use is mainly characterized by pastures and crops (13%)  
323 and coniferous forest (9%). The isotope-inferred indication of groundwater recharge  
324 across the altered coniferous forests may be explained by the devastation produced by  
325 the bark beetle plague ( $\sim 381 \times 10^3$  ha in central Honduras; [Blair et al., 2019](#)), which may  
326 result in less evapotranspiration rates and in a net recharge and runoff increase.

327 Nevertheless, detailed water balance estimations are still needed to clearly underpin the  
328 ecohydrological effects of this forest degradation and the expansion of crop lands  
329 across high elevations in the basin. Groundwater recharge over 1,855 m asl (class e)  
330 covers only 3% of the study area. This area comprises only a 7% of pastures and crops  
331 and is mostly linked to local recharge of springs and boreholes.

## 332 **6. Conclusions**

333 In Honduras complex interactions of socioeconomic and environmental factors  
334 have given rise to a dynamic mosaic of patches of reforestation, deforestation, urban  
335 and agricultural expansion (Kammerbauer and Ardon, 1999; Southworth et al., 2004;  
336 Munroe et al., 2007; Vignola et al., 2015b). Understanding of groundwater vulnerability  
337 to land use changes (i.e., recharge areas and potential contaminant transport) and  
338 climate variability (i.e., rainfall deficits and recharge amount) is critical in Central  
339 Honduras for the upcoming decades.

340 Our results indicated a large dependency of Caribbean-type moisture in rainfall  
341 events in central Honduras, a moisture transport mechanism often disrupted during the  
342 warm ENSO events. The latter results in extreme rainfall deficits, and consequently, in  
343 low levels of drinking water reservoirs and groundwater recharge.

344 Water stable isotope compositions in rainfall, drilled wells, boreholes, and springs  
345 (within the headwaters of the Choluteca River basin) revealed a strong connectivity  
346 between modern precipitation and groundwater reservoirs. Groundwater isotopic values  
347 (during baseflow periods) are within the range of annual mean  $\delta^{18}\text{O}$  rainfall values in the  
348 valley. While groundwater apparent ages are still needed in this basin, these findings  
349 suggest a high potential of contaminants, pathogens, and solutes transport from  
350 agricultural and urban areas to shallow groundwater reservoirs.

351 Isotope-inferred mean recharge elevations range between 813 to 2,130 m asl  
352 with a mean value of  $\sim 1,600$  m asl and corresponded with coniferous forests, crop  
353 lands, and urban areas. It is important to highlight that the influence of rapid infiltration  
354 of isotopic pulses prior and after baseflow periods (i.e., large rainfall events during the  
355 two peaks of the rainy season) limited the performance of inverse elevation calculations  
356 as follows: springs (60%), borehole (77%), and drilled wells (93%). Drilled wells (up to  
357 300 m depth) denote well-mixed water and -in general- are more representative of the

358 mean annual isotopic recharge composition. High elevation and shallow springs and  
359 boreholes are more susceptible to flashy recharge responses, which in turn may bias  
360 recharge elevation estimates based on annual isotopic lapse rates.

361 Our results are intended to enhance the delineation and mapping of critical  
362 recharge areas in the dense populated (198.4 inhabitants /km<sup>2</sup>; [INE, 2021](#)) headwater  
363 basin of the Choluteca River. Conservation enforcement, municipal land use  
364 regulations, and groundwater extraction permits should be based on a strong scientific  
365 background that guarantees water quality and quantity within the large spectrum of  
366 climate events that Central Americas has been experiencing in the last decade.

### 367 **Data availability statement**

368 Our database is available at  
369 <http://www.hydroshare.org/resource/99b6f970a56740be9b7d6204aeca0553> (García-  
370 Santos et al., 2021). CUAHSI's hydrological repository Hydroshare  
371 (<https://www.hydroshare.org/>) is an online platform to share data, models, and code.

### 372 **Credit author statement**

373 SG-S, RS-M, and TP-P conceptualize the sampling strategies. SG-S, MC-E, and  
374 JH-O conducted all field sampling campaigns. Stable isotopes analysis was conducted  
375 by RS-M and SG-S. Isotope data was curated by RS-M. SG-S, RS-M, and JM-E  
376 initiated the manuscript. All co-authors contributed to the final manuscript version.

### 377 **Acknowledgements**

378 Support from the United Nations Environment Programme (UNEP) adaptation  
379 funds through the project "Ecosystem-based adaptation in the forest corridor of Central  
380 Honduras" was fundamental. This study was also partially-supported from the Scientific  
381 Research Office and Honduran Institute of Earth Sciences at UNAH (Tegucigalpa,  
382 Honduras). The Stable Isotopes Research Group and the Water Resources  
383 Management Laboratory at Universidad Nacional (Heredia, Costa Rica) collaborated  
384 with water stable isotopes analysis. The authors thank Pedro Ortiz (SANAA), Kelly  
385 Almendares (IHCIT), and Maynor Ruiz (UNAH) for logistic sampling support and  
386 geological revisions of a previous manuscript version, respectively.

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**Figure 1:** Study area within the upper basin (headwaters) of the Choluteca River (dark green polygon). Elevation gradient is color-coded in a gray-scale (from 541 to 2,335 m asl). Sampled water sources for stable isotopes analysis are color and symbol-coded: springs (cyan triangles), boreholes (orange squares), drilled wells (red circles), and rainfall (purple crosses). La Brea (upper basin; orange cross) and UNAH (lower basin; purple cross) meteorological stations are included for reference. Main streams are represented by blue lines. The inset shows the location of the Choluteca River basin in Honduras and the Central America Isthmus.

**Figure 2:** Monthly rainfall seasonality between the upper (La Brea station; 1,610 m asl; blue bars) and lower (UNAH station; 1,063 m asl; pink bars) basin areas. Long-term (La Brea: 1976-2012; UNAH: 1979-2019) standard deviations are denoted by crossed bars.

**Figure 3:** Major lithological features within the headwaters of the Choluteca River

**Figure 4:** Monthly  $\delta^{18}\text{O}$  (‰) (A) and deuterium excess (‰) (B) variability at three rainfall monitoring stations (Cerro de Hula, El Picacho, and UNAH) across the headwaters of the Choluteca River basin from 2018 to 2020 (daily, N=270 and weekly, N=156). Black arrows represent five main climatic features: E-W trade winds, easterly waves, MSD, tropical cyclones season and ITCZ activation.

**Figure 5:** Dual isotope diagrams including: drilled wells (A), boreholes (B), and springs (C). Rainfall mean values at UNAH (daily), El Picacho and El Cerro de Hula (weekly) are denoted by color-coded circles and bi-directional error bars. Grey crosses represent the rainfall isotopic variability within the Choluteca River basin (CBMWL; gray line). The GMWL (pink line) is included as reference.

**Figure 6:** HYSPLIT air mass back trajectories (48 hours backwards) coded by  $\delta^{18}\text{O}$  (‰) (A) and  $d$ -excess (‰) in daily rainfall at Tegucigalpa, Honduras (black star) (2018-2020; N=270).

**Figure 7:** (A) Map of potential recharge elevations within the headwaters of the Choluteca River basin. The background shows a color-coded digital elevation map from 541 to 2,335 m asl. Color-coded circles denote groundwater samples (springs, N=128; drilled wells, N=166; and boreholes, N=70) with resolved elevations from a basin-wide isotopic lapse rate. Grey circles represent sites where large seasonal variability limited the inverse recharge elevation calculation. (B) Mean recharge elevation histogram.

**Figure 8:** Land use types within the headwaters of the Choluteca River basin in 2019. The Institute of Conservation and Forestry Development (Tegucigalpa, Honduras) defined 19 classifications which denote the high degree of land fragmentation. The most prominent recharge areas cover elevations from 813 to 2,130 m asl with a mean value near 1,600 m asl.

